Structural Geology and Timing of Deformation at Gibraltar Cu-Mo Porphyry Deposit; Cariboo Region, British Columbia

Introduction

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he Gibraltar Cu-Mo porphyry, located northwest of Williams _ake is hosted in the Granite Mountain batholith (217 M \pm sysouth et al., 1995). The main ore zone, hosted within the Vine Series Phase tonalite, is structurally dismembered several different deformation events. Questions still exist regarding the relationship between mineralization (215 \pm 1. -210 ± 0.9 M (Harding, 2012)) and the earliest deformation.

his study is part of the Targeted Geoscience Initiative project which aims to expand knowledge of de intrusion-related ore deposits across Canada.

Objectives

Unravel the geometry and kinematics of deformatio that have affected ore distribution;

Place constraints on the timing of deformation structure

) Determine if batholith emplacement and mineralizatior were syn-kinematic with the earliest deformation structur or if structural modification of the deposit occurred p emplacement and mineralization

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Left: Regional geology with Gibraltar mine in centre inset with four main pits outlined. Inset top left: British Columbia and the Cache Creek and Quesnel terranes and their spatial relationship to Gibraltar mine. Right: Gibraltar pit map with the locations of the 7 bench walls mapped as part of this study. Granite Lake Faults (GLF represented as thrust faults on map) are cross cut by N-S trending oblique-slip faults mapped by Oliver (2007, 2008). Modifications to some N-S trending faults within the Granite Lake pit are based on detailed structural mapping and ore displacement modelled on Leapfrog Geo 3-D visualization tools.

The Granite Mountain batholith contains highly sheared mafic phases south of the mine property and becomes increasingly more felsic towards the north, consisting of diorite, quartz diorite phases, tonalite phase and trondjhemite; locally, intrusion of late leucocratic quartz porphyry dikes are observed.

The Mine phase tonalite is relatively equigranular with an average grain size of 2-4 mm, comprising of ~15-25% qtz, 40-50% plagioclase and 25-35% chlorite (altered hornblende) and hosts most of the mineralization at Gibraltar porphyry.

The deformation intensity generally contains a positive correlation between the alteration and mineralization (e.g. Oliver, 2007; van Straaten et al., 2013). Veins predominantly comprise different alteration zone assemblages. The chart below is based on field mapping and drill core logging observations: this data was used to construct cross sections of the Granite Lake operational pit

Hydrothermal Alteration Assemblage	Alteration Characteristics	Vein Assemblage	Vein Shape and Texture	Mineralization Stage
Saussurite-Chlorite (Albite-Epidote-Zoisite)	No alteration to pale yellow-green saussuritization of feldspars, chloritized-Hbl and presence of epidote veinlets	Ер	1 mm planar veinlets and 4-5 cm wide diffused flooding	Pre- mineralization
Propylitic (Chlorite-Epidote)	Increase in pale yellow-green saussuritization of feldspars, chloritized-Hbl, epidote grains and veinlets, and Chl-Ep veins*	Chl+Ep±Py±C py±Qtz±Cb	1-15 mm wide Chl-Ep vein. a) thin, planar; b) wider, diffuse margins; c) wider, diffuse Qtz- envelope; d) Cb and cubic Py in the centre ±Cpy	Early
Chlorite-Quartz	Alteration intensity characterized by vein density and ranges from no pervasive matrix alteration to prevalent Qtz and Chl replacement of Fsp*	Qtz±Chl±Mag ±Py±Cpy±Mo	2-20 mm wide Qtz-vein with Chl-halo. Sometimes Mag/Chl/Mo/Cpy±Py aligned in centre. a) sharp boundaries; b) no margins, grey Qtz; and c) disconnected, wavy veins with more diffuse Qtz-Chl margins*	Main
Quartz-Sericite	Qtz-Ser flooding	Qtz+Ser	Qtz-Ser flooding and replacement of Chl-Qtz- Fsp alteration	Late
Quartz-Sericite-Chlorite	Finely disseminated Ser± pale Chl*±Qtz alteration of matrix. Euhedral grains of Py are sparse	~	No specific vein is closely associated with this alteration	Late
Phyllic (Quartz-Sericite-Pyrite)	Occurs in varying intensities. Weak QSP alteration is distinguished by 1-3 cm wide sheeted veins, while stronger QSP alteration is characterized by pervasive replacement of the matrix by Qtz and Ser*	Qtz+Ser+Py± Cpy±Mo	a) 1-3 cm wide sheeted grey Qtz-veins, with Ser- Qtz envelopes and cubic Py aligned in the centre. b) 1-200 cm wide milky-white veins, with parallel sheeted Mo-veinlets, host bulk Mo- mineralization (Harding, 2012)	Late
Ankerite-Quartz	Pale Ank-Qtz alteration commonly associated with high strain zones. Sulphide mineralization may occur with Ser±Chl folia*	Ank-whisps	2 mm in size, separated sinuous whisps. Veins were either completely deformed or transposed as they are unidentifiable	Late
~	Not associated with any specific alteration assemblage	Qtz+Chl±Cpy ±Py±Cb	10 cm-1 m thick, boudinaged Qtz-veins with Chl- knots ±Py±Cpy blebs, enveloped by Chl/Ser-folia	Late or Post- mineralization

Alteration assemblages are based on the predominant (>50%) alteration assemblag Foliation intensity increases with decreasing space between form (strike) lines



blique to inclined boudinaged gtz-veir









Modified from van Straaten et al., 2013

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Field Observations



ed Mine phase tonalite with S chloritized hbl, ep and chl. Drill core sample





with ~2m wide cataclastic zone. High-angle fault cross-cuts and drags S1 foliation, displaying

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Relative Timing of Intrusion and Deformation I: Foliation overprints mineralized veins





Right: S1 foliation defined by sericite and a weak chlorite alignment, the sericite foliation overprints previously saussuritized, and veined Mine phase tonalite and refracts into a steeper orientation in the vein.





mostly by elongate qtz.

hl-knots and large py and cpy blebs



The orebody in Granite Lake pit is confined to a "panel" within the two main GLFs (Oliver, 2008), and are offset by NW-SE trending dextral, oblique strike-slip faults. Orebody was interpolated on Leapfrog and projected onto a 2013 pit map with faults mapped by Oliver (2008), and modified by the author using microstructural analysis, field observations and Leapfrog.

II: Sheridan stock

A: Foliated Sheridan stock tonalite with foliation oriented 085/46, looking down on

B: Variably foliated tonalite. Foliation defined by elongate qtz and sericite-illite lamellae. Foliation is E-W striking, and south dipping as observed in locations demarcated by yellow circles on the map. Data from Schiarizza, 2014.

the Gibraltar porphyry





III: Temperature of deformation

Left: S1 defined by elongated qtz sericite lamellae. Folded with vergence towards the NE.

Weak subgrain rotation and bulging recrystallization f atz are indicative of dislocation creep and temperatures of >300°C to <400°C in dry conditions.

However, given the involvement of hydrothermal alteration and veining, fluids likely play a large role in deformation and would ultimately lower the required temperature for gtz deformation.

Plagioclase is relatively undeformed and displays minor undulose extinction indicating <450°C.

Right: Moderately dipping, and crenulated S1, parallel to mineralized vein with cpy. Cpy brittly deformed by shallowly dipping S2 crenulation cleavage.











owing the use of his field data and photographs.